Extreme wave value analysis under uncertainty of climate change scenarios off Iberian Peninsula coast

Bárbara F. V. Vieira*, José L. S. Pinho, Joaquim A. O. Barros

Department of Civil Engineering, School of Engineering, University of Minho, Braga 4710-057, Portugal

*Corresponding author, barbaravasquezvieira@gmail.com

Abstract

Extreme wave value analysis under uncertainty scenarios was developed to estimate wave climate characteristics at 17 stations in southwestern European coast off the Iberian Peninsula. A comprehensive wave dataset downscaled with the Wave Watch III (WWIII) model by Meteogalicia under MarRisk Project was used considering results of models from the Coupled Model Intercomparison Project 5 (CMIP5). Descriptive statistics for significant wave height (Hs), peak wave period (Tp), and mean and peak wave direction were performed for historical data (1960-2005), and for projected data in two twenty-year time periods under two Representative Concentration Pathway (RCP) scenarios (2026–2045 and 2081–2100). Hs and Tp extreme values for the study area were obtained using the Gumbel, Fréchet and Weibull probability distributions for the 10-, 50-, and 100-year return period. Obtained results showed that: historical Hs values decrease from North to South and are higher than those calculated in any of the RCPs future scenarios; mean Tp values appear to be constant in all stations; and means for peak and mean direction have higher frequency of occurrence in Q4 $(270^{\circ}-360^{\circ})$. This study also allowed the computation of Hs and Tp values for 100-year return period, which can be used as design criteria for structural analyses in maritime works.

Keywords

Extreme events, probability distribution, significant wave height, uncertainty, wave climate projections, RCPs scenarios

1. Introduction

The special characteristics of coastal zones including their high population density has greatly increased during the recent decades, which is associated to rapid economic growth and coastward migration (Merkens et al., 2016; Neumann et al., 2015). Demographic pressure on coastal areas has resulted in significant infrastructure and assets being located at risk-prone areas, increasing exposure and vulnerability to natural hazards along the coast.

Indeed, coastal zones are extremely vulnerable regions with delicate physiographic equilibria, whose ecosystems are highly influenced by Mean Sea Level Rise (MSLR) and related hazards (including erosion, flooding and salt intrusion) that are expected to significantly increase by the end of this century in the absence of major additional adaptation efforts (Callaghan et al., 2020; Cheng and Chen, 2017; Chini et al., 2010; Ding et al., 2013; Feng et al., 2019; Mori et al., 2013; SROCC, 2019; Warner and Tissot, 2012; Xie et al., 2019). The evolution over the past two centuries suggests that the tendency for sea levels rising and consequent coastal erosion will aggravate in the 21st century independently of the considered global warming scenarios (Mase et al., 2013; Mori et al., 2013; Mori et al., 2014).

Climate change and its undesirable consequences, such as an expected increasing frequency and magnitude of extreme events, generate additional risks to water-related infrastructure, requiring an ever-increasing need for adaptation measures (SROCC, 2019). Special planning and management approaches with a paradigmatic shift from crisis management to risk management (hazard analysis and vulnerability analysis) in a changing environment are required.

Risk management approaches require actual and future projections of wave climates, including storm wave data. For this purpose, application of probability distributions for extreme wave climate data is a usual applied methodology to estimate extreme wave data for the design, operation and maintenance of coastal infrastructures, under uncertainty caused by climate change.

The United Nations Intergovernmental Panel on Climate Change (IPCC) adopted different greenhouse gas concentration trajectories (Representative Concentration Pathway – RCP), to describe different future climates based on the volume of greenhouse gases emitted in the future years. The RCPs are commonly used to assess scientific, technical and socio-economic information concerning climate change, its potential effects and options for adaptation and mitigation (IPCC, 2020).

Predicted change in global mean surface temperature and global MSLR for the mid- and end 21st century (2046–2065 and 2081–2100 averages, respectively), relative to the 1986–2005 period, is strongly dependent on which RCP emission scenario is followed. MSLR is projected to rise between 0.47 m (0.32–0.63 m, likely range) (RCP4.5) and 0.63 m (0.45–0.82 m, likely range) (RCP8.5) by 2100 (likely range) relative to 1986–2005 (IPCC, 2014). These sea levels rise will imply the dissipation of wave energy at higher levels in the nearshore that could be exacerbated by eventual aggravation of the extreme wave climates.

The most used descriptor of the wave field is the energy-density spectrum in both frequency and direction of propagation. From this spectrum, most of the parameters

commonly used for describing wave climate regimes can be derived, namely: the significant wave height, $H_s(m)$, the peak wave period, $T_p(s)$, the mean wave direction, $D_m(^\circ)$, and the peak wave direction, $D_p(^\circ)$. Three main types of wave data are available: from observation, measurement or simulations. The observation or measurement of waves require personnel and measuring equipment *in situ* at the time of observation, whereas the simulated wave data are produced and operated by many major meteorological services by making use of numerical wave modelling (WMO, 1998) whose results could be used in the design of coastal structures.

The most widely used variables to design offshore and onshore structures are H_s , T_p and D_m (Capitão and Fortes, 2011; Carvalho and Capitão, 1995; Park et al., 2020; WMO, 1998). Much of the effort given to wave climate studies in recent years has concentrated upon statistical methods for estimating extreme values of these parameters, which require representativeness, consistency, and validity of the collected data to avoid incorrect extrapolation. Gumbel, Fréchet, Weibull, and log-normal value distributions are mostly applied to derive extreme wave data (Capitão and Fortes, 2011; Martucci et al., 2010; Mathiesen et al., 1994; Park et al., 2020; Silva et al., 2008; Wang et al., 2013).

In this study, downscaled wave data of MeteoGalicia (Bio et al., 2020; Pinho et al., 2020) has been used for describing three different wave climates recurring to statistical analysis of H_s , T_p , D_m , and D_p and in estimating extreme wave data values at 17 selected stations in the Atlantic Ocean off the Iberian Peninsula. This dataset consists of 46 years of historical data (1960–2005) and two twenty-year time periods of RCPs projected data (2026–2045 and 2081–2100).

Although extreme H_s values computed by wave models could underestimate real values (Dentale et al., 2018; Reale et al., 2020), an extreme value analysis was applied to the datasets defined from the calculation of the 95th percentile of H_s , in order to identify storm waves (Castelle et al., 2015; Goda, 2000; Harley, 2017; Masselink et al., 2014). The 10-, 50-, 100-year return period (T_r) of H_s and T_p were estimated using the Gumbel, Fréchet, and Weibull distributions. Performance studies applying these three methods concluded that Fréchet and Weibull distributions fit better in estimating extreme values of H_s (Capitão and Fortes, 2011; Carvalho and Capitão, 1995; Guedes-Soares and Carvalho, 2001; Mathiesen et al., 1994; Piccinini, 2006; Vanem, 2015). Projections from RCP4.5 and RCP8.5 datasets also allow to introduce uncertainty in the results obtained, since the effects of plausible climate change scenarios were considered.

Previous studies in the western coast of Portugal were based on limited time series periods and considered single locations over the region (e.g., Capitão and Fortes, 2011, and Silva et al., 2008 using a time series of 14 years collected at a buoy located at Leixões; and Guedes-Soares and Carvalho, 2001 using a time series of 1 year, collected at a buoy located off port of Sines).

Following those studies, more updated research on extreme value analysis and descriptive statistics on the variability of wave parameters off the Iberian Peninsula, especially at the Portuguese coast, is needed to better understand the wave climate at this region. Based on this recognized gap, this research work intends to address a comprehensive analysis on the wave extremes variability off the Iberian Peninsula under climate change scenarios.

The comprehensiveness of the wave climate regime database including historical data and projected data for two different greenhouse gases emission RCP4.5 and RCP8.5 scenarios, as well as the statistical methods used, allowed a descriptive statistics and an extreme value analysis of the local wave climate regimes.

Outcomes of this research work include the estimation of extreme values of H_s and T_p . This information can be a valuable contribution for the establishment of design parameters for coastal engineering projects in the study area and for coastal vulnerability analysis of the Atlantic Iberian coast to climate change.

2. Study Area and Data Sources

The Iberian Peninsula (Portugal and Spain) has an Atlantic coastline with approximately 1.300 km of extension. This energetic coast has been subject to intensive urbanisation and coastward migration associated mainly to tourism, fishing, aquaculture, industry, and port activities. In the last decades, this region has being subject to an erosional process, coastal floods and aquifers salinization as showed in different studies (Álvarez et al., 2020; Baptista et al., 2014; Cherneva et al., 2005; Guedes-Soares and Carvalho, 2001; Pereira and Coelho, 2013; Pinho et al., 2020).

Winter storms are of common occurrence in this region (Gomes et al., 2018; Santos et al., 2018). Due to scarcity of monitoring data, hydrodynamic modelling works have been performed to assess the effect of three categories of storms on water levels at this coast. Ensemble simulations performed by Gomes et al. (2018) estimated a maximum sea level of 2.3 m for hurricanes and tropical storms, and a maximum sea level of 1.2 m for tropical depressions (values above mean spring-tide level of 3.75 m). The effects of these phenomena combined with storm waves can be associated with wave-overtopping and coastal flooding events with harmful consequences for coastal environments.

In this study the wave climate data (significant wave height, peak period, and wave direction) were obtained from wave datasets downscaled with the Wave Watch III model (WWIII) of Meteogalicia, both for historical and future climates. This model was forced with results of the Model for Interdisciplinary Research On Climate 5 (MIROC5), which is included in the Coupled Model Intercomparison Project 5 (CMIP5) atmospheric/ocean global simulations (Qu et al., 2020; Sperna-Weiland et al., 2016; Vanem, 2017). HP original CMIP5 (MIROC5) results were obtained by properly assimilate ocean and atmospheric observed data as reproduced in Tatebe et al. (2012).

Within the scope of the MarRisk project (Bio et al., 2020; Pinho et al., 2020), the Meteogalicia WWIII model was forced with CMIP5 data, allowing this way to downscale model results for 17 stations off the western coast of Iberian Peninsula. Figure 1 presents the study area and the location of those stations.



Figure 1. Study area with location of the 17 stations around the Iberian Peninsula.

The spectral data resulting from the WWIII model runs was provided for the 17 selected stations, which are properly numbered and aggregated according their location (North, West, and South) in relation to the Iberian Peninsula (Table 1).

Station	Longitude	Latitude	Depth related to mean sea level (m)
N1	-8.00	44.00	-206,60
N2	-9.00	44.00	-1199,45
N3	-10.00	44.00	-4971,26
WN1	-11.00	44.00	-4998,02
WN2	-11.00	43.00	-2490,85
WN3	-11.00	42.00	-2734,10
WN4	-11.00	41.00	-3890,95
WS1	-11.00	40.00	-4723,50
WS2	-11.00	39.00	-3333,70
WS3	-11.00	38.00	-5018,40
WS4	-11.00	37.00	-2533,70
WS5	-11.00	36.00	-4870,75
S1	-10.00	36.00	-4542,95
S2	-9.00	36.00	-3700,45
S 3	-8.00	36.00	-1522,45
S4	-7.00	36.00	-809,55
S 5	-6.00	36.00	-161,93

 Table 1. Coordinates of the study selected stations.

WWIII spectral data results comprises historical wave climate data, as well as wave climate projections based on the RCPs climate change scenarios. The CMIP5 selected model used to simulate climate change scenarios was possible, since it have yielded reasonable calibration results using a sound 46 years wave climate data.

Results obtained from the application of recently developed phase 6 of CMIP (CMIP6) demonstrated similar outcomes as those obtained with CMIP5 (McKenna et al., 2020; Plecha and Soares, 2020; Zhu et al., 2020). While there was a significant improvement

from CMIP3 to CMIP5 in simulating mean sea level, the same was not verified from CMIP5 to CMIP6. Major differences are mainly found in middle-to-high latitudes in the Southern Hemisphere. Moreover, regarding ocean dynamic sea level projections, a similarity of results was also verified between CMIP5 and CMIP6 (Lyu et al., 2020).

The used data includes five different wave climate scenarios: one for the historical period, HP, (1960–2005), and four projections under the climate change scenarios RCP4.5_mid (2026–2045), RCP4.5_end (2081-2100), RCP8.5_mid (2026–2045), and RCP8.5_end (2081-2100). The historical period includes 46 years of data, while the projection periods comprise two twenty-year time periods (for the mid and end periods of 21st century) which give a reasonably extensive datasets enabling reliable extrapolations for wave climate regimes analysis (Capitão and Fortes, 2011; Park et al., 2020; Wang et al., 2016; Wang and Swail, 2001).

The WWIII spectral information files, in NetCDF format, are divided into 1512 files making up 12.9 GB of disk space. Each NetCDF file contains spectral information for one month and the temporal resolution of the data is 3 hours. Table 2 shows the structure of the information used in each of the NetCDF files.

Variable	Description	Туре	Unit
direction	Sea surface wave direction	1D	degree
Dpt	Depth	2D	m
Efth	Sea surface wave directional variance spectral density	2D	m ² s rad ⁻¹
frequency	Frequency of centre band	1D	s ⁻¹
frequency1	Frequency of lower band	1D	s ⁻¹
frequency2	Frequency of upper band	1D	s ⁻¹
latitude	Latitude	2D	Degree North
longitude	Longitude	2D	Degree East
station	Station id	1D	-
station_name	Station name	-	-
time	Julian day	1D	day

Table 2. List of variables used in each WWIII NetCDF file.

3. Methodology

3.1. Methods for extreme wave data analysis under uncertainty

Waves extremes eventually exacerbated by climate change increase the magnitude of hazards to coastal infrastructure, either expressed in terms of frequency or severity (IPCC, 2014; Marone et al., 2017; Pinho et al., 2020). The use of probability models is of paramount importance for determining wave climate regimes and predicting intensity and duration of extreme events occurrences. Climate projections for the 21st century are inherently uncertain since there is low confidence in projections of many aspects of climate phenomena that influence regional climate change (Sun et al., 2017). The rate of future global warming effects depends on unpredictable natural influences on climate like volcanic eruptions, as well as on greenhouse gas concentrations due to anthropogenic activities (IPCC, 2020).

Uncertainties arising from climate change scenarios may question the validity of future projections based on data collected in the past. Some authors claim that "stationarity is dead" stating that climate change undermines basic assumption that historically has facilitated management of water resources and risks (Milly et al., 2008). However, this philosophical approach has been contradicted by many authors who consider that the process of climate change is accepted as a 'certainty' and value the need for hydrological predictions based on assumptions that should include stationarity (Montanari and Koutsoyiannis, 2014). In line with these authors, this research work follows a pragmatic approach in predicting extreme wave climate regimes based on the exploitation of simulated datasets generated by numerical models applied to historical and to RCPs scenarios.

An extensive temporal and spatial wave database was used to effectively deal with climate change uncertainty. The usual parameter chosen to estimate the severest conditions likely to be experienced by coastal infrastructures is either the 50- or the 100-year return period of wave height (H_s), where the N-year return period is defined as that which is exceeded on average once every N years (WMO, 1998). This parameter is very important in coastal vulnerabilities analysis and coastal defence solutions design. However, for a complete description of extreme conditions, information on wave peak period (T_p) is also required. This research work has concentrated upon methods for estimating extreme values of H_s, and the correspondent T_p.

Techniques for extreme value analysis consist of adjusting a theoretical probability distribution function to the function of the estimated distribution of a sample, in order to describe a certain random variable behaviour (Silva et al., 2008). Several methods for estimating extreme wave data and to define the local wave climates are available elsewhere (Karian and Dudewicz, 2000; Naghettini and Pinto, 2007; Nascimento, 2009; Park et al., 2020; Sansigolo, 2008; Thevasiyani and Perera, 2014; Urošev et al., 2016).

For this purpose, different probability distributions have been applied: Gumbel (Cotta et al., 2016; Silva et al., 2008); Gumbel and Weibull (Mathiesen et al., 1994; Park et al., 2020); Gumbel, Fréchet and Weibull (Capitão and Fortes, 2011; Carvalho and Capitão, 1995; Goda, 2000; Sansigolo, 2008).

In order to assess the relative performance of different methods, this research applied three generalized extreme value (GEV) probability distributions (Gumbel, Fréchet, and Weibull) to estimate the extreme values of the H_s and the associated T_p for the 10-, 50-, 100-year return period. The application of these probability distributions required a previous selection of annual maxima H_s values for all stations. Usually, two different methods (annual maxima and peaks-over-threshold) are widely used in extreme value analysis. According to Goda (2000), the annual maxima approach should be used for databases covering more than 20 years, while for shorter records of extreme data the peaks-over-threshold seems to be a more reliable method. For methodological coherence reasons, in this study the annual maxima method was selected for the 46-year span HP data and for the two 20-year span RCPs scenarios data. These GEV probability distribution methods consider that the distribution of maxima of *n* values are asymptotic with increasing the number *n*, and this is the reason for their usual application on extreme meteorological events with good results.

Extreme value analysis was conducted after transforming the extreme value distributions from the probability distribution function form to the cumulative distribution function (CDF) form for the convenience of computation. Equations (1) to (3) show the CDF of Gumbel, Fréchet and Weibull distributions (Vivekanandan, 2012):

$$F(X) = e^{-e^{-\left(\frac{X-\alpha_G}{\beta_G}\right)}}, \beta_G > 0 \text{ and } -\infty < X < \infty \text{ (for Gumbel)}$$
(1)

$$F(X) = e^{-\left(\frac{X}{\beta_F}\right)^{(-\lambda_F)}}, \, \beta_F > 0 \text{ and } -\infty < X < \infty \text{ (for Fréchet)}$$
(2)

$$F(X) = e^{-\left(\frac{X}{\beta_W}\right)^{\lambda_W}}, \beta_W > 0 \text{ and } -\infty < X < \infty \text{ (for Weibull)}$$
(3)

where *X* is the random variable under consideration (H_s, T_p); α_G and β_G are the location and scale parameters of the Gumbel distribution; β_F , λ_F and β_W , λ_W are the scale and shape parameters of Fréchet and Weibull distributions, respectively.

The inverse first-order reliability method (I-FORM), as proposed by Winterstein et al. (1993), was applied to calculate the exceedance probability according to the Equation (4):

$$P = S(X_i) = i/(N+1), \ P \in]0,1[$$
(4)

where *P* is the exceedance probability of *X*; *N* is the total number of samples; *i* (i=1,...,N) is the number of sample values for which any variable value $x \le X_i$.

The method of fitting the chosen distribution was based on the application of the probability plot correlation coefficient technique (Filliben, 1975). The use of this graphical technique requires the computation of the percent point function, which is the inverse of the cumulative distribution function. The formulae for the percent point function G(P) expressing graphical coordinates that fit the Gumbel, Fréchet, and Weibull distributions are presented in Table 3.

	Percent Point Function Plotting				
Probability Distribution	x-axis coordinate	y-axis coordinate			
Gumbel	X _i	$G_g(P) = -\ln[\ln(1/P)]$			
Fréchet	$\ln(X_i)$	$G_f(P) = -\ln[-\ln(P)]$			
Weibull	$\ln(X_i)$	$G_w(P) = \ln[ln[1/(1-P)]]$			

 Table 3. Percent point functions for the selected probability distributions.

Equation 5 was employed to estimate the 10-, 50-, 100-year return period of extreme wave data, which gives the correspondence between the exceedance probability, P, and the return period, T_r (Silva et al., 2008).

$$T_r = \frac{1}{1-P} \tag{5}$$

The chosen distribution function is then fitted by approximating a linear function through the plotted values by the least squares method. The parameters were obtained using the slopes and y-intercept values when the coefficient of determination (R^2) value was the highest.

3.2. Descriptive statistics of the local wave regimes

A comprehensive statistical analysis was performed to determine the characteristics of wave regimes for each of the stations located off the Iberian Peninsula coastal zone. The downscaled CMIP5 results were used in the construction of the following five wave climate regimes datasets: (i) HP; (ii) RCP4.5_mid and RCP4.5_end; and RCP8.5_mid and RCP8.5_end. A diagrammatic scheme that shows the followed descriptive statistics analysis is presented in Figure 2.

A statistical analysis was performed for: (i) the complete wave dataset; and (ii) for a storm sub-dataset, obtained from the application of a criterion based on H_s values exceeding the 0.95 quantile as storm definition (a peaks-over-threshold method where a set of values above a certain threshold level are selected). This methodology was applied to obtain characteristics of wave climate evolution with special focus on storms due to the relevance of extreme phenomena.

Wave dataset served as the basis for determining the mean values for H_s , T_p , D_m , and D_p (independent variables); the maximum values for H_s ; and the 0.95 quantile of H_s ($H_{s,0.95}$). For the purpose of this study, a storm event was considered when the significant wave height exceeds $H_{s,0.95}$ (Castelle et al., 2015; Harley, 2017, Masselink et al., 2014). From the calculation of $H_{s,0.95}$, a sub-dataset was obtained with mean storm values. Based on this sub-dataset, descriptive statistics were obtained for peak H_s , T_p , D_m , and D_p . Moreover, for the storm event data, a frequency analysis was also conducted to analyse which D_m and D_p quadrants are more frequent at each station. An analysis of the maximum, and mean storms duration, as well as the number of storms registered and the mean storms per year at each station was also performed.



Figure 2. Methodological scheme applied for descriptive statistics.

3.3. Extreme value analysis of wave climate regimes

The extreme value analysis of storm sub-dataset was performed to estimate the 10-, 50-, 100-year return period of H_s and T_p for all stations off the Iberian Peninsula coastal

zone, covering the five wave climates regimes. The method employed followed two sequential steps: (i) the wave regime historical data for the period of 1960-2005 was selected and the Gumbel, Fréchet, and Weibull GEV probability distributions were used comparing which of them give better fitting of extreme wave regime values, by calculating the respective coefficient of determination (R^2) ; (ii) selected GEV probability distributions (Fréchet for H_s, and Weibull for T_p, based on the best fitting method) were applied to estimate the extreme values for two time periods (2026-2045 and 2081–2100) under the RCP4.5 and RCP8.5 climate change scenarios. Figure 3 depicts the scheme adopted in estimating the extreme values for wave regimes.



Figure 3. Methodological scheme applied for extreme value analysis.

4. Results and discussion

At southwestern coast of Iberian Peninsula variable wave climatic regimes are notorious depending on the different climate scenarios and on the geographical location. Descriptive statistics for complete wave dataset and storm sub-dataset, as well as extreme value analysis of Hs and Tp (for historical and projected data) have been performed for all analysed stations. Obtained results are presented in a graphical form supported by tables that include wave climate data numerical values. Comparison with existing studies and proposal of extreme values for coastal engineering design are also discussed.

The stations were clustered into four groups as presented in Table 4. Geo-location of the stations and the homogeneity of the obtained results served as criteria to define the station groups: N for North stations; WN and WS for West stations; and S for South stations.

Table	Table 4. Station groups.						
Station Group	Stations						
N	N1; N2; N3						
WN	WN1; WN2; WN3; WN4						
WS	WS1; WS2; WS3; WS4; WS5						
S	S1; S2; S3; S4; S5						

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4.1. Descriptive statistics: complete wave dataset

Main descriptive statistical results for mean H_s , $H_{s,0.95}$, maximum H_s , mean T_p , mean D_m and mean D_p values for wave dataset are presented in a graphical form in Figure 4.



Figure 4. Mean H_s, H_{s,0.95}, maximum H_s, mean T_p, mean D_m and mean D_p values for wave dataset under five wave climate regime scenarios.

Results for wave dataset are organized in Table 5, where minimum and maximum values for each station group for the five wave climate regimes are presented.

Wave	Station	Wave climate regime scenario							
data	Group	HP	RCP4.5_mid	RCP4.5_end	RCP8.5_mid	RCP8.5_end			
	Ν	2.52-2.87	2.45-2.80	2.33-2.69	2.40-2.75	2.27-2.63			
Mean H _s	WN	2.84-2.97	2.79-2.91	2.69-2.80	2.76-2.85	2.65-2.73			
(m)	WS	2.68-2.81	2.66-2.77	2.56-2.67	2.63-2.74	2.54-2.63			
-	S	1.25-2.58	1.23-2.56	1.19-2.45	1.22-2.52	1.18-2.43			
	Ν	5.20-5.78	5.07-5.62	4.84-5.42	5.00-5.63	4.91-5.52			
II	WN	5.48-5.91	5.30-5.73	5.07-5.55	5.34-5.72	5.27-5.62			
H s,0.95 (III)	WS	4.92-5.38	4.78-5.20	4.54-4.94	4.78-5.22	4.68-5.17			
-	S	2.82-4.82	2.73-4.66	2.58 - 4.40	2.69-4.67	2,69-4,56			
-	Ν	13.98-15.17	15.70-16.77	12.66-14.88	13.19-14.86	15.79-20.35			
Max. H _s	WN	14.70-15.41	13.99-16.79	14.63-16.17	11.36-14.90	15.81-20.74			
(m)	WS	13.82-15.59	10.66-13.37	10.91-15.63	10.42-11.28	12.22-13.95			
(111)	S	8.54-12.58	8.32-10.51	7.97-10.77	7.87-10.33	9.22-12.32			
-	Ν	10.61-10.89	10.45-10.73	10.14-10.47	10.37-10.65	10.07-10.39			
Mean T _p	WN	10.57-10.67	10.40-10.50	10.10-10.23	10.32-10.41	10.00-10.08			
(s)	WS	10.69-10.73	10.53-10.58	10.27-10.35	10.44-10.53	10.11-10.20			
	S	10.05-10.92	9.88-10.78	9.67-10.58	9.89-10.76	9.57-10.44			
-	Ν	278-290	275-289	269-286	276-289	274-289			
Mean D _m	WN	273-299	271-298	264-297	271-300	269-297			
(°)	WS	305-314	307-316	308-317	308-317	307-316			
	S	274-313	272-314	271-314	274-314	271-314			
	Ν	269-285	265-283	259-279	263-281	264-282			
Mean D _p	WN	252-258	245-252	239-247	245-253	247-256			
(°)	WS	267-275	262-272	261-270	263-271	270-277			
	S	277-311	274-311	274-312	276-312	274-312			

Table 5. Descriptive statistics results for the local wave regime using wave dataset from CMIP5.

Significant wave height values (mean H_s , $H_{s,0.95}$, and maximum H_s) decrease, in general, from the northern to southern station groups in the sequence WN>N>WS>S, where values found for S are considerably lower than those of the remaining station groups. Comparing the HP results for mean H_s , $H_{s,0.95}$, and maximum H_s among the station groups, it was found that $H_{s,0.95}$ values are twice the mean H_s values, and the maximum H_s is thrice the $H_{s,0.95}$ values (ranges from 1.25–2.97 m for mean H_s ; 2.82–5.91 m for $H_{s,0.95}$; and 8.54–15.59 m for maximum H_s).

For each station group, when historical values are compared with the projected results of the RCPs scenarios for mean H_s and $H_{s,0.95}$, tendencies show that HP results are higher than those calculated in any of the RCPs scenarios. For example, differences in percentage between RCP8.5_mid scenario and HP for the most energetic station group (WN) show H_s minimum and maximum values of: -2.82 % and -4.04 % for mean H_s ; -2.55 % and -3.21 % for $H_{s,0.95}$; and -22.72 % and -3.31 % for max H_s . Regarding the differences between RCP8.5_end scenario and HP, the corresponding percentages are: -6.69 % and -8.08 % for mean H_s ; -3.83 % and -4.91 % for $H_{s,0.95}$; and +7.55 % and +34.59 % for max H_s . Similar results for the Iberian coast can also be found in the assessment of CMIP5 wave projections conducted by Morim et al. (2019). These results are very important as they contradict results obtained by other authors when the effects of climate change on H_s are simulated. Wang and Swail (2001) showed significant

linear increasing trends on wave height in both the North Atlantic and North Pacific, when the global model hindcast results were analysed for the time period of 1958–1997. In a more recent study, Wang et al. (2016) presented results showing H_s increases in the tropics (especially in the eastern tropical Pacific) and in southern hemisphere high-latitudes, which are based on the CMIP5 historical (2005), RCP4.5 (2050) and RCP8.5 (2099) forcing scenario simulations. Moreover, tendencies show that projected values obtained for RCP4.5 and RCP8.5 scenarios have slight differences apparently without major influence on these wave statistical results.

Maximum H_s values show contradictory tendencies when HP data are compared to projected data. The maximum values obtained for RCP4.5_mid and RCP8.5_end scenarios are higher (10–35 %) than those of the HP data for the northern station groups N and WN (ranging from 16.79 m to 20.74 m, respectively) and lower for the southern station groups WS (13.37 m to 13.95 m) and S (10.51 m to 12.32 m). For RCP4.5_end and RCP8.5_mid scenarios, the results obtained follow the same tendency as verified with mean H_s , and $H_{s,0.95}$, with lower values than those of the HP data. This can be justified by the fact that the use of model data often originates undervaluation of the H_s in projected scenarios.

Mean T_p appears to maintain constant values around 10.50 s in all station groups, independently the scenario considered.

Wave direction defined by D_m and D_p values show slight differences both in station group location and in wave climate regime scenarios. However, it can be noted that for D_m the values for N and WN station groups have analogous behaviour with maximum values around 300°, and in WS and S station groups maximum values around 315°. D_p maximum values around 280°, 250°, 275°, and 312° can be found in station groups N, WN, WS, S, respectively.

4.2. Descriptive statistics: storm sub-dataset

Storm sub-dataset for the five wave climate regime scenarios were used to calculate mean peak storm H_s , mean storm T_p , total number of storms, mean storms per year and storm durations. Figure 5 presents the results obtained for these five wave climate regime scenarios.

Table 6 shows minimum and maximum values obtained for storm wave data in each station group for the five wave climate regime scenarios. Mean peak storm H_s and mean storm T_p , as well as total number of storms, mean storms per year, mean storm duration, and maximum storm duration are presented.

For mean peak storm H_s , differences in percentage between RCP8.5_mid scenario and HP for the most energetic station group (WN) show H_s minimum and maximum values of +1.01 % and -1.21 %. Regarding the differences between RCP8.5_end scenario and HP, the corresponding percentages are +4.49 % and +3.23 %.

Wave	Station		Wave o	limate regime s	scenario	
data	Group	HP	RCP4.5_mid	RCP4.5_end	RCP8.5_mid	RCP8.5_end
Mean	Ν	6.59-7.28	6.60-7.26	6.53-7.24	6.57-7.14	6.71-7.50
peak	WN	6.90-7.42	6.96-7.49	6.98-7.43	6.97-7.33	7.21-7.66
storm H _s	WS	6.09-6.73	6.13-6.77	6.09-6.85	6.11-6.77	6.29-6.88
(m)	S	3.77-5.98	3.78-5.97	3.74-5.98	3.71-5.99	3.91-6.17
	Ν	14.02-14.28	13.83-14.15	13.54-13.95	13.66-14.15	13.83-14.14
Mean -	WN	13.94-14.03	13.81-14.11	13.44-13.75	13.54-13.66	13.72-13.83
storm I _p	WS	13.93-14.00	13.75-14.03	13.66-13.83	13.57-13.77	13.88-14.05
(8)	S	12.52-14.04	12.56-13.80	12.29-13.89	12.38-13.80	13.04-14.06
T-4-1	Ν	819-827	314-324	291-296	301-327	281-282
1 otal	WN	798-820	281-301	253-289	311-327	272-282
of storms	WS	754-805	278-288	228-248	283-303	244-269
	S	648-747	256-284	206-226	258-283	222-241
	Ν	18-18	16-16	15-15	15-16	14-14
Mean	WN	17-18	14-15	13-14	16-16	14-14
storms per vear	WS	16-18	14-14	11-12	14-15	12-13
per jeur	S	14-16	13-14	10-11	13-14	11-12
Mean	Ν	21-22	21-22	19-19	21-22	21-23
storm	WN	22-22	22-24	19-22	21-22	22-23
duration	WS	22-24	23-25	21-23	22-24	24-26
(h)	S	24-28	23-28	22-27	23-26	26-31
Max.	Ν	168-174	132-135	93-123	117-171	117-126
storm	WN	177-198	132-135	123-135	108-117	123-141
duration	WS	132-138	108-135	135-162	87-99	138-228
(h)	S	147-243	135-201	162-204	102-156	144-228

Table 6. Descriptive statistics results for the local wave regime using storm sub-dataset (1).



Figure 5. Mean peak storm H_s , mean storm T_p , total number of storms, mean storms per year and storm durations results under five wave climate regime scenarios.

Mean peak storm H_s decrease, in general, from the northern to southern groups of stations in the sequence WN>N>WS>S, where values found for S are considerably lower (approximately, one-half in minimum values) than those of the remaining station groups. The results obtained for RCPs scenarios are similar to the historical ones, notwithstanding there is a slight increase in these values for the climate scenario RCP8.5_end.

The results obtained for mean storm T_p show similar values in the range [13–14 s], independently the station groups and the scenario considered.

Total number of storms values decrease in the sequence of N>WN>WS>S, as well as from the historical (maximum values ranging from 747–827 in 46 years) to the RCPs (maximum values ranging from 226–327 in 46 years) scenarios. Mean storms per year have the same behaviour as total number of storms, with values ranging from 10 to 18 events per year.

Results for mean storm duration values show a slight increase in the sequence of N>WN>WS>S (with maxima in the range of 22–28 h), while small variations for the different scenarios are verified. Maximum storm duration values have a decreasing behaviour in the sequence of N>WN>WS>S for HP data, and an increase behaviour for RCP4.5_end and RCP8.5_end in the same sequence. Maximum storm duration projected values for RCPs in S are higher than in N, WN, and WS with maximum values ranging from 156 h to 228 h.

Figure 6 presents mean D_m and mean D_p results obtained from the storm sub-dataset for the five wave climate regime scenarios. Frequency analysis of mean D_m and mean D_p , associated to peak storm wave height for the fourth quadrant (Q4: 270°–360°) are also presented in Figure 6.



Figure 6. Results for mean D_m, mean D_p, and frequency analysis of mean D_m and mean D_p for Q4.

Table 7 shows minimum and maximum values obtained for mean D_m and D_p , in each station group for the five wave climate regime scenarios. In addition, for these two wave data, the range of percentages of their occurrence in the Q4 [270°–360°] related to the other three quadrants are also presented.

Wave	Station		Wave climate regime scenario					
data	Group	HP	RCP4.5_mid	RCP4.5_end	RCP8.5_mid	RCP8.5_end		
	Ν	282-296	283-298	281-294	282-296	281-291		
Mean D _m	WN	278-294	279-296	278-292	280-289	276-281		
(°)	WS	297-305	297-309	294-306	297-309	290-299		
_	S	271-303	272-305	271-303	272-306	273-298		
	Ν	284-297	287-300	285-296	285-298	282-290		
Mean D _p	WN	277-286	276-281	275-284	283-286	279-282		
(°)	WS	290-298	277-282	288-296	287-294	286-297		
_	S	275-306	275-306	275-302	276-310	277-302		
O4 Ema a	Ν	68-86	67-85	68-86	67-88	65-81		
Q4 Freq	WN	67-78	65-79	66-72	67-73	61-65		
(9)	WS	80-89	78-91	73-89	80-91	73-87		
(70)	S	63-88	65-90	59-87	64-90	68-86		
	Ν	77–92	78-91	73-89	78–95	71-86		
Q4 Freq.	WN	75-85	75-86	70-80	76-83	68-75		
(%)	WS	86-95	86-94	86-93	88-95	83-94		
()	S	81-94	82-94	79-93	82-96	85-95		

Table 7. Descriptive statistics results for the local wave regime using storm sub-dataset (2).

The results obtained for either mean D_m or mean D_p show similar values in the range [271°–310°], independently the station groups and the scenario considered. For HP data, the occurrence of the wave directions falling in the Q4 have a frequency ranging from 63 % to 89 % for mean D_m and from 75 % to 94 % for mean D_p . It is also relevant to mention that the second most frequent wave direction falls in the third quadrant (Q3: 180°–270°).

4.3. Extreme values analysis

4.3.1. Historical data

Storm data was used as the basis for extreme wave climate regime analysis. Since the probability distribution methods (Gumbel, Fréchet and Weibull) applied in this study are built on specific assumptions, a comparison of the estimates obtained by each of these methods has been performed. The three methods used the HP dataset for estimating H_s and T_p of three return periods: 10-, 50-, 100-year. As a representative example, Figure 7 shows the results obtained for determining extreme values for H_s and T_p in one of the stations, applying the percent point function plotting technique.



Figure 7. Parameter estimation of extreme value distribution using the least squares method for H_s and T_p in station WN3.

Table 8 summaries the estimates for wave data (H_s and T_p) applying the three probability distribution methods in all stations for the considered return periods. The obtained coefficient of determination (R^2) value, which served as selection factor of the most appropriate method to apply in the RCP4.5 and RCP8.5 scenarios, is presented in Table 9. From the calculated results, it can be noted that Fréchet distribution overestimates wave heights for longer return periods. However, this distribution provided higher R^2 values in estimating H_s results, whereas Weibull distribution has demonstrated to give the best fit (higher R^2 values) for the T_p .

	-	H _s (m)			$T_{p}(s)$			
Station	Tr (years)	Gumbel	Fréchet	Weibull	Gumbel	Fréchet	Weibull	
	10	12.00	12.19	11.64	18.56	18.72	18.13	
N1	50	14.07	15.12	12.46	21.04	21.91	19.05	
	100	14.95	16.57	12.72	22.08	23.41	19.35	
	10	12.72	12.85	12.35	18.24	18.40	17.82	
N2	50	14.90	15.83	13.21	20.64	21.49	18.72	
	100	15.82	17.28	13.49	21.66	22.94	19.00	
	10	13,01	13,12	12,63	18,17	18,29	17,73	
N3	50	15,23	16,12	13,51	20,67	21,51	18,68	
•	100	16,16	17,58	13,79	21,73	23,03	18,99	
	10	13.23	13.38	12.85	17.96	18.12	17.53	
WN1	50	15.52	16.53	13.76	20.44	21.34	18.47	
•	100	16.49	18.08	14.05	21.49	22.86	18.77	
	10	12.93	13.04	12.56	18.18	18.41	17.72	
WN2	50	15.18	16.08	13.45	20.80	21.91	18.71	
	100	16.13	17.57	13.74	21.91	23.58	19.03	
	10	12.63	12.68	12.27	18.71	19.01	18.17	
WN3	50	14.84	15.59	13.14	21.56	22.96	19.25	
• • • •	100	15.77	17.01	13.42	22.77	24.87	19.60	
	10	12.34	12.36	11.98	18.69	19.02	18.18	
WN4	50	14.56	15.28	12.88	21.44	22.83	19.20	
•	100	15.50	16.71	13.16	22.60	24.66	19.53	
	10	12.14	12.10	11.78	18.62	18.93	18.12	
WS1	50	14.36	14.94	12.68	21.36	22.71	19.15	
-	100	15.31	16.34	12.97	22.52	24.53	19.48	
	10	11.86	11.73	11.47	18.25	18.44	17.80	
WS2	50	14.08	14.45	12.36	20.78	21.76	18.76	
-	100	15.02	15.79	12.64	21.86	23.34	19.06	
	10	11.62	11.43	11.21	18.04	18.22	17.60	
WS3	50	13.86	14.10	12.09	20.46	21.38	18.51	
	100	14.81	15.41	12.38	21.48	22.87	18.80	
	10	11.30	11.10	10.92	18.40	18.66	17.92	
WS4	50	13.46	13.66	11.78	21.26	22.59	19.03	
	100	14.38	14.91	12.06	22.47	24.50	19.39	
	10	10.88	10.79	10.55	18.25	18.59	17.75	
WS5	50	12.88	13.27	11.35	21.20	22.79	18.90	
	100	13.73	14.48	11.61	22.44	24.83	19.27	
	10	10.47	10.40	10.20	18.59	18.98	18.06	
S1	50	12.28	12.64	10.93	21.61	23.33	19.24	
	100	13.05	13.73	11.17	22.89	25.46	19.61	
	10	10.11	10.06	9.85	18.73	19.10	18.20	
S2	50	11.86	12.24	10.56	21.69	23.32	19.33	
	100	12.60	13.30	10.78	22.94	25.38	19.70	
	10	9.43	9.42	9.15	18.28	18.69	17.73	
S3	50	11.14	11.63	9.84	21.48	23.39	18.98	
	100	11.86	12.72	10.06	22.83	25.72	19.39	
	10	8.69	8.70	8.40	17.54	17.92	16.99	
S4	50	10.45	11.08	9.12	20.67	22.51	18.22	
	100	11.19	12.28	9.35	21.99	24.79	18.61	
	10	7.50	7.56	7.22	17.09	17.41	16.56	
S 5	50	9.13	9.93	7.91	20.41	22.30	17.92	
	100	9.82	11.14	8.13	21.81	24.75	18.36	

Table 8. Extreme wave climate analysis for the historical scenario in all stations. Comparison of H_s and T_p values for 10-, 50-, and 100-year return periods obtained with the application of Gumbel, Fréchet and Weibull probability distributions.

Station		H _s (m)			T _p (s)	
Station	Gumbel	Fréchet	Weibull	Gumbel	Fréchet	Weibull
N1	0.954	0.922	0.961	0.958	0.936	0.957
N2	0.960	0.941	0.931	0.959	0.934	0.960
N3	0.970	0.953	0.929	0.968	0.951	0.942
WN1	0.979	0.955	0.943	0.963	0.939	0.954
WN2	0.977	0.961	0.921	0.949	0.918	0.968
WN3	0.989	0.982	0.907	0.909	0.879	0.974
WN4	0.993	0.993	0.886	0.908	0.870	0.990
WS1	0.973	0.986	0.842	0.928	0.889	0.989
WS2	0.940	0.973	0.796	0.961	0.935	0.969
WS3	0.906	0.955	0.758	0.961	0.932	0.975
WS4	0.901	0.952	0.733	0.937	0.908	0.943
WS5	0.958	0.984	0.808	0.943	0.900	0.976
S1	0.963	0.983	0.797	0.921	0.878	0.976
S2	0.961	0.977	0.804	0.915	0.873	0.980
S 3	0.969	0.976	0.856	0.926	0.887	0.979
S4	0.962	0.964	0.859	0.920	0.888	0.974
S 5	0.973	0.963	0.905	0.936	0.913	0.928

Table 9. Extreme wave climate analysis for the historical scenario in all stations. Comparison of R^2 values obtained with the application of Gumbel, Fréchet and Weibull probability distributions.

4.3.2. Projected data

Extreme wave climate regime projections for RCP4.5 and RCP8.5 scenarios were performed using the Fréchet and the Weibull methods in estimating H_s and T_p , respectively. Wave data values were calculated for 10-, 50-, 100-year return periods.

Table 10 summarizes the estimation of wave data under the RCPs scenarios for 10-, 50-, and 100-year return periods at all stations.

			DCD8 5 mid						
Station	Tr (years)	$\frac{\mathbf{KCP4}}{\mathbf{W}}$	<u>5_mia</u>		.5_ena		.5_mia		.s_ena
	10	$H_{s}(m)$	T _p (S)	$H_{s}(\mathbf{m})$	Tp (S)	$H_{s}(m)$	Tp (S)	$H_{s}(m)$	T _p (S)
N 14	10	13.36	17.91	11.80	17.04	12.12	18.01	13.23	17.62
NI	50	18.28	18.97	14.44	17.87	15.71	18.91	17.63	18.61
	100	20.87	19.31	15.73	18.13	17.53	19.20	19.90	18.93
	10	14.17	17.79	13.10	16.90	12.87	18.02	15.37	17.29
N2	50	18.96	18.97	16.62	17.81	16.76	19.18	21.70	18.07
	100	21.44	19.35	18.38	18.11	18.74	19.56	25.11	18.32
	10	14.62	18.01	13.66	17.38	13.18	17.62	15.96	17.23
N3	50	19.42	19.20	17.47	18.50	17.05	18.77	22.64	17.95
	100	21.90	19.58	19.39	18.86	19.01	19.14	26.25	18.18
	10	14.67	18.20	13.93	16.65	13.32	17.77	16.32	17.19
WN1	50	19.47	19.36	17.74	17.41	17.26	18.94	23.14	17.87
	100	21.94	19.73	19.66	17.66	19.26	19.32	26.81	18.09
	10	13.87	17.90	14.15	16.99	12.94	17.49	15.80	17.63
WN2	50	18.05	19.16	18.71	17.77	16.39	18.62	22.12	18.49
	100	20.17	19.56	21.05	18.02	18.11	18.98	25.50	18.76
	10	13.00	18.06	13.98	17.18	12.27	17.39	15.10	17.33
WN3	50	16.55	19.39	18.80	18.00	15.12	18.45	20.76	18.13
	100	18.33	19.82	21.31	18.26	16.51	18.79	23.75	18.39
	10	12.42	18.01	13.49	17.11	13.18	17.62	15.96	17.23
WN4	50	15.76	19.22	18.04	17.90	17.05	18.77	22.64	17.95
	100	17.44	19.61	20.40	18.15	19.01	19.14	26.25	18.18
	10	12.19	18.76	13.12	17.16	10.80	17.13	13.68	17.20
WS1	50	15.69	20.30	17.69	18.01	12.34	18.15	18.19	17.89
	100	17.46	20.80	20.07	18.28	13.06	18.48	20.53	18.11
	10	11.70	18.55	12.62	17.02	10.58	16.74	13.22	17.20
WS2	50	14.89	20.02	17.11	17.79	11.98	17.60	17.69	17.93
	100	16.49	20.50	19.45	18.03	12.63	17.87	20.01	18.16
	10	11.43	18.61	11.87	17.26	10.58	16.70	12.63	16.97
WS3	50	14.60	20.21	15.80	18.20	12.13	17.56	16.64	17.74
	100	16.18	20.73	17.83	18.49	12.85	17.83	18.70	17.98
	10	11.20	17.36	11.31	17.24	10.41	16.76	12.46	17.39
WS4	50	14.10	18.72	14.94	18.24	12.13	17.74	16.76	18.19
	100	15.54	19.15	16.81	18.55	12.93	18.05	19.00	18.45
	10	11.18	17.83	10.75	17.12	9.92	16.74	12.13	17.10
WS5	50	14.35	19.28	14.05	18.22	11.51	17.72	16.20	17.86
	100	15.94	19.75	15.73	18.57	12.25	18.03	18.31	18.10
	10	11.09	17.94	10.46	17.64	9.73	16.82	11.80	17.22
S1	50	14.47	19.54	13.63	18.94	11.28	17.79	15.64	18.16
	100	16.20	20.06	15.24	19.36	12.00	18.10	17.62	18.46
	10	10.96	18.13	10.13	17.96	9.56	16.85	11.52	17.57
S 2	50	14.58	19.68	13.14	19.21	11.29	17.79	15.36	18.53
5-	100	16.45	20.18	14.67	19.61	12.12	18.09	17.34	18.83
	10	10.21	15.77	9.54	17.52	8.69	17.14	10.78	17.17
S 3	50	13.72	16.47	12.58	18.94	10.31	18.17	14.58	18.12
~ -	100	15.55	16.69	14.14	19.40	11.09	18.50	16.57	18.42
	10	9.03	15.57	8.87	17.17	7,83	16.93	9,74	17.33
S4	50	12.19	16.37	12.06	18.51	9.50	18.07	13.46	18.50
54	100	13.84	16.62	13 73	18.95	10.31	18.43	15 44	18.88
	10	7.64	15.29	7.89	17.35	6.85	17.00	8.45	17.27
85	50	10.40	16.01	11 27	19.22	8.66	18 70	12.06	18 59
60	100	11.85	16.24	13.11	19.84	9.56	19.75	14.02	19.01
	100	11.05	10.24	1.5.11	17.04	1.50	17.40	17.04	17.01

Table 10. Extreme wave climate values estimated by using the Fréchet and Weibull probabilitydistributions under the RCP4.5_mid, RCP4.5_end, RCP8.5_mid and RCP8.5_end scenarios for 10-, 50-,
and 100-year return periods at all stations.

4.3.3. Significant wave height

Results obtained for different H_s return periods (Table 11) show increasing values from 10- to 100-year, as expected. In general, the values of H_s decrease from the northern to southern station groups in the sequence WN>N>WS>S, where values found for S are considerably lower (average from 67 % to 77 %) than those of the remaining station groups.

For each station group, when HP data values are compared with the projected results from the RCPs scenarios for H_s, tendencies show that: (i) HP data values are slightly analogous than those calculated in the RCPs_mid; (ii) projected values obtained from RCPs_end have higher values when compared to HP data values, with special incidence of the results obtained from RCP8.5.

H_s return values obtained from RCP4.5_mid and RCP4.5_end are apparently contradictory, since results are mostly lower for RCP4.5_end than for RCP4.5_mid. However, similar results can also be found in Copernicus Climate Change Service (2019) for the Iberian Peninsula. This is not the case with the RCP8.5 scenarios, where the return values from RCP8.5_end are consistently higher than those from RCP8.5_mid.

Return	Station	Wave climate regime scenario						
(years)	Group	HP	RCP4.5_mid	RCP4.5_end	RCP8.5_mid	RCP8.5_end		
	Ν	12.19-13.12	13.36-14.62	11.80-13.66	12.12-13.18	13.23-15.96		
10	WN	12.36-13.38	12.42-14.67	13.49-14.15	12.27-13.32	15.10-16.32		
10	WS	10.79-12.10	11.18-12.19	10.75-13.12	9.92-10.80	12.13-13.68		
	S	7.56-10.40	7.64-11.09	7.89-10.46	6.85-9.73	8.45-11.80		
	Ν	15.12-16.12	18.28-19.42	14.44-17.47	15.71-17.05	17.63-22.64		
50	WN	15.28-16.53	15.76-19.47	17.74-18.80	15.12-17.26	20.76-23.14		
50	WS	13.27-14.94	14.10-15.69	14.05-17.69	11.51-12.34	16.20-18.19		
	S	9.93-12.64	10.40-14.58	11.27-13.63	8.66-11.29	12.06-15.64		
	Ν	16.57-17.58	20.87-21.90	15.73-19.39	17.53-19.01	19.90-26.25		
100	WN	16.71-18.08	17.44-21.94	19.66-21.31	16.51-19.26	23.75-26.81		
100	WS	14.48-16.34	15.54-17.46	15.73-20.07	12.25-13.06	18.31-20.53		
	S	11.14-13.73	11.85-16.45	13.11-15.24	9.56-12.12	14.02-17.62		

Table 11. Return values for H_s (m).

4.3.4. Wave peak period

For each return period, wave peak period T_p (Table 12) appears to maintain constant values in all station groups, independently the scenario considered. Approximate values of 17,50 s, 18,00 s, and 19,00 s have been obtained for 10-, 50-, 100-year return period, respectively.

Return	Station	Wave climate regime scenario						
(years)	Group	HP	RCP4.5_mid	RCP4.5_end	RCP8.5_mid	RCP8.5_end		
	Ν	17.73-18.13	17.79-18.01	16.90-17.38	17.62-18.02	17.23-17.62		
10	WN	17.53-18.18	17.90-18.20	16.65-17.18	17.39-17.77	17.19-17.63		
10	WS	17.60-18.12	17.36-18.76	17.02-17.26	16.70-17.13	16.97-17.39		
	S	16.56-18.20	15.29-18.13	17.17-17.96	16.82-17.14	17.17-17.57		
	Ν	18.68-19.05	18.97-19.20	17.81-18.50	18.77-19.18	17.95-18.61		
50	WN	18.47-19.25	19.16-19.39	17.41-18.00	18.45-18.94	17.87-18.49		
50	WS	18.51-19.15	18.72-20.30	17.79-18.24	17.56-18.15	17.74-18.19		
	S	17.92-19.33	16.01-19.68	18.51-19.22	17.79-18.70	18.12-18.59		
	Ν	18.99-19.35	19.31-19.58	18.11-18.86	19.14-19.56	18.18-18.93		
100	WN	18.77-19.60	19.56-19.82	17.66-18.26	18.79-19.32	18.09-18.76		
100	WS	18.80-19.48	19.15-20.80	18.03-18.57	17.83-18.48	17.98-18.45		
	S	18.36-19.70	16.24-20.18	18.95-19.84	18.09-19.25	18.42-19.01		

Table 12. Return values for T_p (s).

4.4. Comparison with existing studies

Two studies (Carvalho and Capitão, 1995 and Capitão and Fortes, 2011) were carried out to estimate 10-, 50-, 100-year return period of H_s at three sites on the western coast of Portugal (Leixões, Figueira da Foz and Sines). These studies differ partly in their research approach and scope from the present one, but are anyway useful benchmark to test the validity of the obtained results.

The main important difference is that the present study uses an extensive modelled dataset consisting of 46 years of historical data and two twenty-year time periods of RCPs projected data applied at 17 locations far from the coast, whereas in the other studies the wave data have been recorded for 13 years in the harbours of Figueira da Foz and Sines (Carvalho and Capitão, 1995) and 14 years in Leixões (Capitão and Fortes, 2011).

Table 13 compares the H_s extremes values predictions of this study for the HP scenario against those obtained in the other two studies. Stations WN3, WS1 and WS3 were considered geographically near to Leixões, Figueira da Foz and Sines, respectively. The comparison between the results obtained in the present study and in Carvalho and Capitão (1995) shows small differences. However, large discrepancies were found when the comparison is made with Capitão and Fortes (2011). The discrepancies found can be justified for three main reasons: (i) the temporal and spatial dimensions of data sources; (ii) the fact that a storm sub-dataset extracted from the wave dataset was used in the calculation of extreme values in the present study; and (iii) the wave climate data near the coast (as used by those authors) is influenced by wave refraction due to local bathymetric characteristics.

		Extreme values of H _s (m)			
Location	Tr (years)	Present study (HP data)	(Carvalho and Capitão, 1995)	(Capitão and Fortes, 2011)	
WN3 Leixões	10	12.69	_	7.25	
	50	15.59	_	9.62	
	100	17.01	-	10.84	
WS1 Figueira da Foz	10	12.14	10.60	_	
	50	14.36	12.80	_	
	100	15.31	13.80	_	
WS3 Sines	10	11.43	10.70	_	
	50	14.10	15.80	_	
	100	15.41	18.20	-	

Table 13. Estimate values of H_s on the western coast of Portugal obtained in different studies.

4.5. Extreme values for coastal engineering design

The results obtained in this study for 100-year return values of H_s and T_p derived for HP data and RCP8.5_end in four station groups are summarized in Table 14. The methodology followed for extreme value analysis based on a storm sub-dataset obtained from the H_s 95th percentile of original wave datasets justifies the higher estimated values for 100-year H_s from the RCP8.5_end scenario when compared to the one estimated for HP data. The obtained results are aligned with other similar works (e.g., Copernicus Climate Change Service (2019)) that project more intense extreme H_s for the end of the century.

Awareness on the prediction of 100-year return values from RCP scenarios should be taken, since these scenarios are based on uncertainty and a 20-year span data. However, the obtained results are robust and can be of valuable interest in engineering practice.

These values can be used as design parameters in maritime structures projects to be applied in the Iberian Peninsula coastal zone, once propagated to the local of interest by appropriate methodologies.

Wave data		Station Group			
		Ν	WN	WS	S
HP data	H _s (m)	16.57-17.58	16.71-18.08	14.48-16.34	11.14-13.73
	T _p (s)	18.99-19.35	18.77-19.60	18.80-19.48	18.36-19.70
RCP8.5_end	H _s (m)	19.90-26.25	23.75-26.81	18.31-20.53	14.02-17.62
	$T_{p}(s)$	18.18-18.93	18.09-18.76	17.98-18.45	18.42-19.01

Table 14. 100-year return period values for H_s and T_p as parameters for structural engineering design.

5. Conclusions

The present study developed an extreme wave value analysis under uncertainty scenarios using comprehensive simulated wave datasets downscaled by Meteogalicia from CMIP5 models to 17 stations off the Iberian Peninsula coastal zone.

The extreme wave data values for H_s and T_p were calculated through the Gumbel, Fréchet and Weibull distributions for the 10-, 50-, and 100-year return period. Five wave climate scenarios using HP data and projected data of RCP4.5_mid, RCP4.5_end, RCP8.5_mid and RCP8.5_end were considered. Descriptive statistical analysis of the obtained wave results demonstrated that: (i) H_s has different values at different locations in the coastal zone with decreasing values from North to South; (ii) HP data values are higher than those calculated in any of the RCPs scenarios; (iii) projected H_s values obtained from RCP4.5 and RCP8.5 have slight differences (approximately 4 %) between these considered scenarios; (iv) maximum H_s values show contradictory tendencies when HP data are compared to projected data; (v) maximum values of H_s obtained for RCP4.5_mid and RCP8.5_end scenarios are higher than those of the HP data for the northern stations and lower for the southern stations; (vi) mean T_p appears to maintain constant values in all stations, independently the wave climate regime scenario considered; (vii) results for mean D_m and mean D_p show higher frequency of occurrence in Q4 (270°–360°) followed by a smaller frequency in Q3 (180°–270°).

Extreme H_s values for all return periods have different values in different locations in the coastal zone with decreasing values from North to South, being considerably lower on the most southern stations. In each station, HP data values are similar to those calculated in the RCPs_mid, and projected values obtained from the RCPs_end have higher values when compared to HP data values. For each return period, wave peak period T_p results show constant values in all stations, independently the scenario considered.

Wave extreme value analysis for the H_s and T_p is an essential element for marine structural design. The high-quality of the database applied and the assumptions taken in this research work allowed the determination of 100-year return period of H_s and T_p values that can be used with confidence as design values for structural analyses in maritime works to be built in the western coast off Iberian Peninsula.

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Author contributions

Bárbara Vieira: Conceptualization, Methodology, Validation, Formal analysis, Investigation, Resources, Writing—original draft preparation, Writing—review and editing, Visualization.

José Pinho: Conceptualization, Methodology, Validation, Formal analysis, Investigation, Resources, Writing—review and editing, Visualization, Supervision.

Joaquim Barros: Conceptualization, Methodology, Validation, Formal analysis, Writing—review and editing, Visualization, Supervision.

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